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Mitigation of atmospheric perturbations and solid Earth movements in a TerraSAR-X time-series

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DOI: <https://doi.org/10.1007/s00190-011-0515-6>

Posted at the Zurich Open Repository and Archive, University of Zurich

ZORA URL: <https://doi.org/10.5167/uzh-50466>

Journal Article

Published Version

Originally published at:

Schubert, A; Jehle, M; Small, David; Meier, E (2012). Mitigation of atmospheric perturbations and solid Earth movements in a TerraSAR-X time-series. *Journal of Geodesy*, 86(4):257-270.

DOI: <https://doi.org/10.1007/s00190-011-0515-6>

Mitigation of atmospheric perturbations and solid Earth movements in a TerraSAR-X time-series

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Received: 12 May 2011 / Accepted: 5 September 2011 / Published online: 24 September 2011
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Abstract The TerraSAR-X (TSX) synthetic aperture radar (SAR) marks the recent emergence of a new generation of spaceborne radar sensors that can for the first time lay claim to localization accuracies in the sub-meter range. The TSX platform's extremely high orbital stability and the sensor's hardware timing accuracy combine to enable direct measurements of atmospheric refraction and solid Earth movements. By modeling these effects for individual TSX acquisitions, absolute pixel geolocation accuracy on the order of several centimeters can be achieved without need for even a single tiepoint. A 16-month time series of images was obtained over a fixed test site, making it possible to validate both an atmospheric refraction and a solid Earth tide model, while at the same time establishing the instrument's long-term stability. These related goals were achieved by placing trihedral corner reflectors (CRs) at the test site and estimating their phase centers with centimeter-level accuracy using differential GPS (DGPS). Oriented in pairs toward a given satellite track, the CRs could be seen as bright "points" in the images, providing a geometric reference set. SAR images from the high-resolution spotlight (HS) mode were obtained in alternating ascending and descending orbit configurations. The highest-resolution products were selected for their small sample dimensions, as positions can be more precisely determined. Based on the delivered product annotations, the CR image

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Keywords Atmospheric path delay · Solid Earth tides · Geolocation accuracy · GPS · Plate tectonics · SAR geometry · TerraSAR-X

1 Introduction

A primary goal of any Earth-observing, remote sensing system is the ability to assign accurate geographic positions to the surface features being imaged. The degree to which it is able to do this is referred to as its *geolocation accuracy*. Until recently, image products from spaceborne synthetic aperture radar (SAR) sensors have all required time-consuming geometric pre-processing, usually involving operator intervention and the collection of ground control points (GCPs) and image tiepoints, before they could be layered with other information sources within geographic information systems. A SAR sensor with high geolocation accuracy greatly simplifies the task of combining multiple data takes with one another, not only simplifying their inter-comparison, but also dramatically speeding up applications such as near-real-time disaster mapping. Accurate geolocation also permits multiple

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image products to be quickly combined or layered with other data sources such as digital elevation models (DEMs), cadastral maps, vegetation maps, forest maps, hydrological maps, etc. Finally, the high geolocation accuracy of TSX raises the possibility of its use as a geometric reference for other instruments (including optical sensors), where the sensor model is not as well known a priori. For example, this could be possible through automatic feature matching based on one or more TSX images and an optical image dataset to be rectified; this technique was already demonstrated by Reinartz et al. (2011) to be feasible.

Recently, the spaceborne SAR system TerraSAR-X (TSX) has begun to provide data with unprecedented geolocation accuracy. TSX is a German Earth observation satellite, built in the context of a public–private partnership between the German Federal Ministry of Education and Research (*Bundesministerium für Bildung und Forschung*; BMBF), the German Aerospace Center (*Deutsches Zentrum für Luft- und Raumfahrt*; DLR) and Astrium GmbH. The launch took place on 15 June 2007, and the sensor has a nominal lifetime of 5 years, although it is currently expected to last up to 7 years. The payload of the satellite is an X-band SAR system with a 9.65 GHz center frequency, an electronically steerable phased-array antenna and a side-looking imaging capability within an off-nadir pivoting range of approximately 20° – 55° . The satellite is in a near-polar dawn/dusk orbit at an altitude of 514 km. Using its active radar antenna, it is able to produce image data with a spatial resolution on the order of ~ 1 m, regardless of weather conditions, cloud cover or absence of daylight.

A SAR sensor mounted on an air- or spaceborne platform measures target positions along two dimensions: its distance from the sensor, or *range* and its along-track position or *azimuth*. These two dimensions are inherent to a SAR measurement and are reflected in the format of the image products generated, called *slant-range* images. The ranging and azimuth timing measurements are independent of each other and thus contribute separately to the sensor's total 3-D geolocation accuracy. While the high geolocation accuracy of the TSX products has already been partially demonstrated (Eineder et al. 2011; Schubert et al. 2010; Weydahl and Eldhuset 2010; Ager and Bresnahan 2009; Jehle et al. 2008; Nonaka et al. 2008; Schubert et al. 2008), all previous studies were partially hindered by one or more uncertainties that have been eliminated in the present study. Eineder et al. (2011) demonstrated that absolute *ranging* accuracy is better than 10 cm for a correctly calibrated instrument, and they measured relative accuracies as high as ~ 3 – 6 cm. However, the along-track (azimuth) geolocation accuracy was not measured in the study. The remaining studies were hindered either by artifacts related to the GPS reference frame, measurement uncertainties or insufficient accounting for one or more perturbations acting upon the SAR signals, such as the

solid Earth tide (SET). Finally, the long-term *stability* of the sensor's localization accuracy had not yet been established firmly; this is treated in this study.

The sensor's high product geolocation accuracy can only be achieved by modeling and correcting for the phenomena that measurably disturb the observed target positions. The refraction of electromagnetic waves in the atmosphere is the largest of these effects, typically causing range delays between 2 and 3 m at X-band. SET perturbations, caused by lunar and solar tidal forces acting on the planet as a whole, are the second-largest error source. They can cause vertical displacements of up to half a meter (peak-to-peak) at the equator (Melchior 1974; Milbert 2011), but typically no more than ~ 40 cm at mid-latitudes. An even smaller effect also exists: Earth deformation by ocean tide loading. As the oceans are periodically redistributed by the relative positions of the sun–moon–Earth system, the changing weight distribution of the water mass on the ocean bed causes the Earth to deform. These vertical fluctuations usually do not exceed several centimeters for inland sites (Penna et al. 2008) and are neglected in this study.

While the atmospheric and solid Earth perturbations cannot be *separated* in the SAR signal, modeling the effects individually for the particular acquisition geometry and test site permits indirect model validation. The problem can be stated as follows: *if* it is possible to accurately predict the location of a reference target in a SAR image having an accurately known position, then the imaging process has been reproduced acceptably in the model-based corrections. As atmospheric and solid Earth movements influence the imaging process, an accurate prediction implies that our understanding of the *processes* involved is equally accurate. Conversely, incorrectly modeling the perturbations will prevent accurate geolocation predictions. This study demonstrates the accuracy and stability of the TSX system, as well as the validity of the atmospheric and physical models used, to within several centimeters of remaining uncertainty.

2 Experimental design

The primary goal of this study was to jointly validate the long-term stability of the TSX sensor and the Earth-system models used to adjust the SAR geolocation measurements. The test site was selected to meet the following criteria: (1) the corner reflectors (CRs) placed there would have to remain stable and undisturbed for the duration of the 16-month test period (security), (2) it would need to be near permanent meteorological and permanent GPS stations and (3) a concrete site would be preferred, to ensure a relatively dark background for the imaged reflectors, permitting accurate estimation of their phase center locations with subpixel accuracy. A test site was selected in the west of Switzerland near the village

Fig. 1 Aerial view of test site Torny-Le-Grand, Switzerland



of Torny-Le-Grand, which met all of the above criteria. Four trihedral CRs were placed at the site, with two permanently facing the TSX sensor perpendicular to its ascending-orbit and two facing its descending orbit. The site was imaged using TSX's highest-resolution spotlight (HS) mode; this time-series of images provided the primary data basis for this study.

2.1 Test site and corner reflectors

The test site is nestled in Switzerland's hilly midlands close to the city of Fribourg. Just outside of the village of Torny-Le-Grand, a former missile launch site of the Swiss Air Force served as an ideal location to place trihedral CRs for the duration of the study. Four CRs were placed about 30 m apart from each other, each surrounded on all sides by enough concrete to guarantee good contrast in the SAR images. The site is shown in Fig. 1, with the four round lots indicated with a yellow circle; a CR was placed in the center of each lot.

A permanent GPS reference station of the Automated GNSS Network of Switzerland (AGNES), located <5 km from the site near the town of Payerne, was used for the differential positioning. Trimble 5700 and Trimble R7 GPS receivers (Fig. 2a) were used to measure their locations with accuracies better than ~1 cm, as given by the RMS reported by the receivers (this accuracy is confirmed by Trimble 2011).

The measurements were repeated midway through and at the end of the campaign to validate the positions used. Existing mounting holes were used as the survey points. Each CR was then mounted such that its phase center (vertex) was in vertical alignment with the survey point (Fig. 2b). The vertical offset between the vertex and survey point was measured, providing an estimated overall absolute accuracy of ~1–2 cm, taking into account all the measurement error sources.

Two of the CRs are seen in Fig. 2c, oriented to face the satellite pass perpendicular to its orbit (i.e. towards the nearest approach). The two western-most CRs were oriented towards the ascending orbit with an elevation angle of 44°; the two eastern-most faced the descending orbit with an elevation of 65°. The ascending and descending geometries corresponded to "long" (~710 km) and "short" (~561 km) paths through the atmosphere, respectively. Their appearance in an HS product (acquired on 25.08.2010) is shown in Fig. 2d. The CR sides have a length of 90 cm. A CR's vertex is its phase center, the point from which the maximum amount of energy is reflected. Its location is given by the peak intensity in the center of the "cross" pattern in a radar image. The cross pattern (actually a *sinc*-like function in the range and azimuth dimensions, caused by constructive and destructive interference of the coherent waves) is the characteristic point target response for a SAR (see section 2.8 in Cumming and Wong 2005). More information on the use of CRs

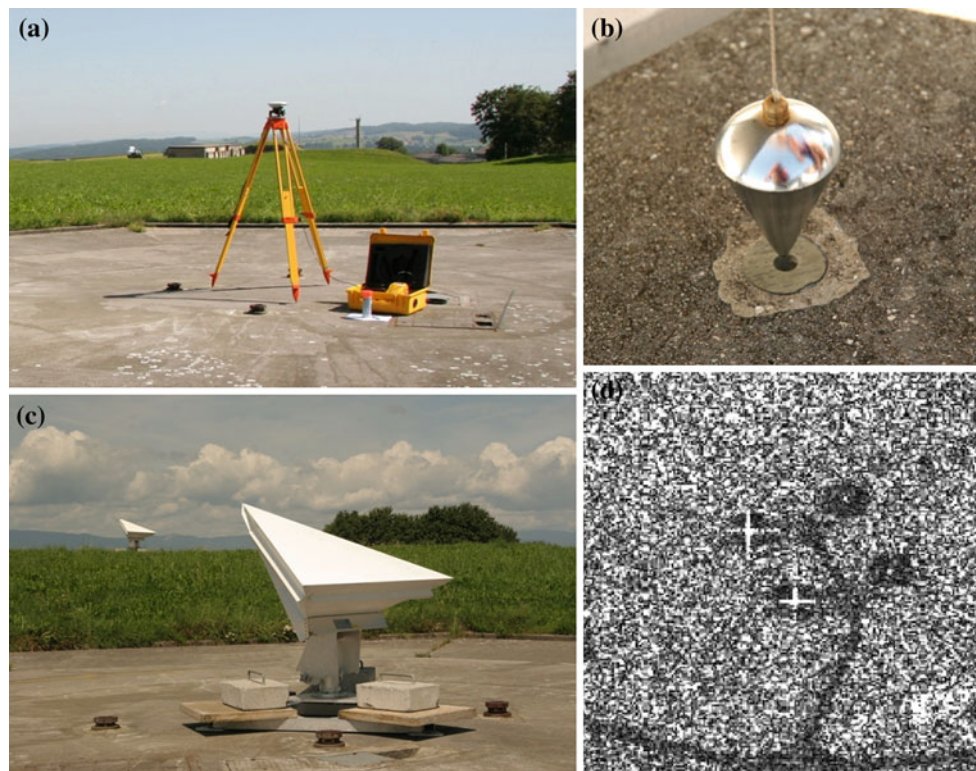
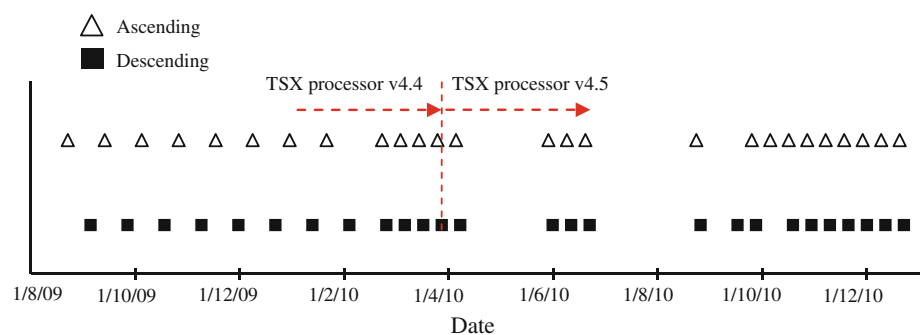


Fig. 2 Surveying positions of the CRs. **a** GPS receiver over a mounting hole, **b** vertex alignment over the mounting hole, **c** CRs oriented towards the descending orbit, and **d** CRs as they appear in an ascending HS product (range spacing 45 cm, azimuth spacing 87 cm; acquisition date 25.08.2010)

Fig. 3 TSX acquisition time series over Torny-Le-Grand test site



as calibration targets is available in section 10-3.3 of [Ulaby et al. \(1982\)](#).

2.2 Time series

The satellite acquires image lines as it advances along its orbit, sending radar echoes down to Earth to the right, relative to the platform's path. A limited number of beams, corresponding to specific elevation angles or look directions, can be selected for imaging. This configuration led to a choice of one particular geometry for ascending orbits and another for descending orbits. The theoretical interval between two

identically configured (repeat-pass) acquisitions is 11 days. The observation period for this experiment was from August 2009 to December 2010. Within this time frame, 52 acquisitions were obtained in total. Of these, 26 were from ascending orbits and 26 from descending orbits. The product acquisition dates are shown graphically in Fig. 3. The series represents a relatively continuous set of observations in a given mode over a fixed test site. Note that at the end of March 2010, the TSX processor was updated with new radiometric and geometric calibration constants and methodologies. The 27 March 2010 acquisition was the first to be processed using the new version of the software. The effect of the recalibration was

clearly evident in our geolocation experiment, manifesting itself as a distinct improvement in the absolute geolocation accuracy.

2.3 Comparison between imaged and measured corner reflector positions

At the core of the test for geolocation accuracy is a comparison between the surveyed location of a CR and its measured location in a SAR image. Given the radar timing annotations (time of first range sample, range sampling rate, first azimuth time, azimuth sample interval) and the state vectors describing the satellite's trajectory during the time of data acquisition, a point on the Earth's surface may be located within the image by solving the Doppler equation governing the image product's geometry. This involves searching for the azimuth time where the satellite's position corresponds to the required Doppler value. This geolocation method is called "range-Doppler" (Meier et al. 1993).

TSX's HS images, with their sub-meter sample spacing, provide a good opportunity to precisely locate reflector peak returns. Reflectors were placed on concrete surfaces to minimize clutter and allow precise location estimates. Applying complex FFT oversampling with a factor of 50 within a 128×128 analysis window centered on a given reflector, its location is determined as the local intensity maximum with a precision of 1/50th of a sample (Small et al. 2004a). The slant range from the sensor to the reflector can then be obtained using the range timing annotations.

The surveyed locations of the CRs were predicted within the radar image products. Additional corrections were made for perturbations described in the following sections, as well as a constant azimuth timing shift indicated in the product annotations for the geolocation grid. The latter provided in the TSX product annotations (Fritz 2007), claims to take "relativistic Doppler" and "internal timing effects" into account. The predicted coordinates were then compared with the corresponding measured positions, yielding a single estimate of the geolocation accuracy. This is illustrated in Fig. 4 for the two reflectors facing the ascending orbit. For each CR, its *predicted* location is marked with a blue cross, the result of transforming the differential GPS (DGPS) coordinate of the CR phase center into slant-range image coordinates. The zoom level increases towards the bottom.

During the estimation of the imaged CR position, a quality indicator is generated at the same time, the signal-to-clutter ratio SCR (Small et al. 2004b). This is defined as the ratio between the CR peak intensity and the mean surrounding intensity (excluding the "arms of the cross" caused by the side-lobes of the CR response). In practice, the mean background intensity is calculated for the four quadrants created by the "cross" pattern, but well outside of the "arms" of the cross. As the SCR is a linear ratio, the radiometric calibration

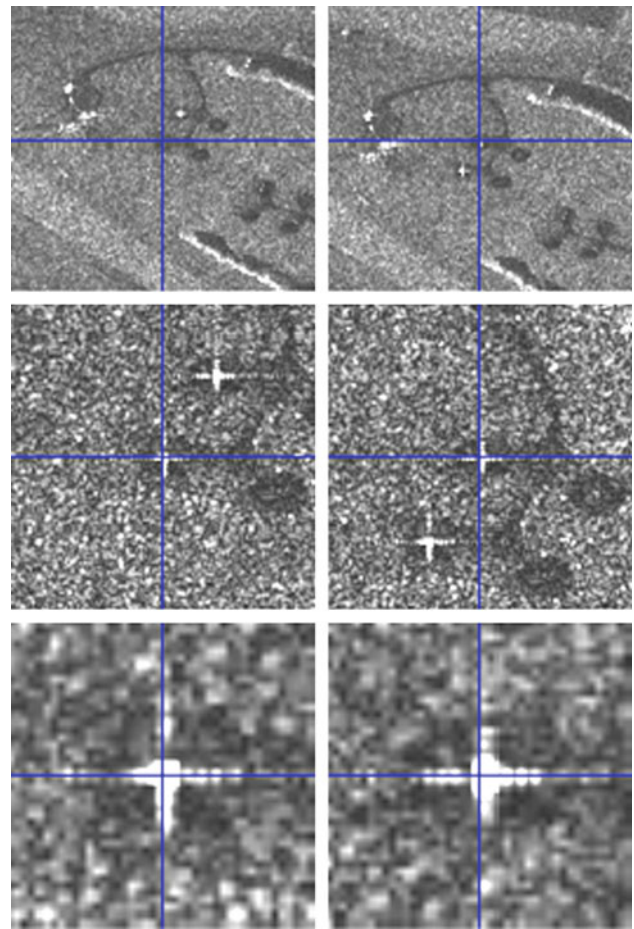


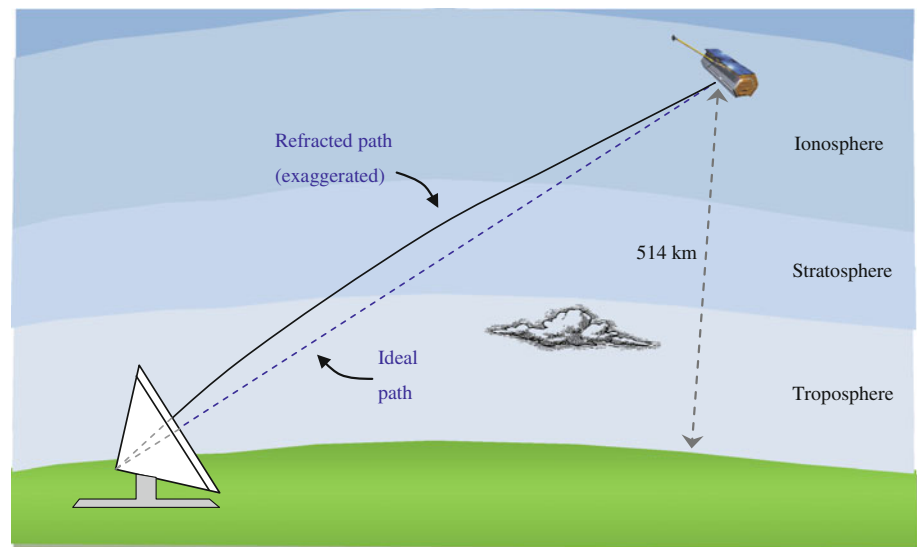
Fig. 4 Both ascending-orbit corner reflectors as seen in a slant-range image, at three zoom levels. The *crosshairs* represent predicted locations based on corner reflector DGPS measurements. Acquisition date 16.08.2009

of the samples is irrelevant. The SCR can be used to automate detection of unreliable measurements, as will be seen later.

3 Perturbations

In the *absence* of an atmosphere, tectonic plate movements and tidal effects, position measurements of stable targets on the Earth's surface using SAR would be predictable and virtually unchanging from day to day. In reality, the sub-metric geolocation accuracy of TSX can only be achieved if proper account is taken of the effects acting on the radar echoes, which alter their measured arrival times at sensor level. The largest contributor to variable deviations in echo return time is atmospheric refraction. The SET cannot be neglected either, but it becomes critical only when an accuracy of 1–2 dm or better is required. Finally, although it does not influence the geolocation accuracy of TSX per se, *estimates* of its accuracy using CRs and GPS may be significantly influenced by plate tectonics due to the use of different

Fig. 5 Atmospheric path delay (atmospheric layers not to scale, refracted path exaggerated)



reference frames. These perturbations were all modeled and removed from our measurements; we were able to show that their removal resulted in very good agreement between theory and measurements. The effects are described in more detail in the following sections.

3.1 Atmospheric path delay

The TSX imaging geometry leads to a path length from sensor to ground of between ~ 550 – 730 km, traversing much of the ionosphere and the entire stratosphere and troposphere. The radar echoes travel more slowly in a medium than in a vacuum (this is why refraction occurs), contributing significantly to the total travel time. This increase in the “ideal” travel time is called the *path delay* (PD). The situation is represented in Fig. 5; the slowing effect of the atmosphere causes the radar echoes to deviate slightly from a straight path, additionally increasing the total traveled distance and hence traveled time. However, the additional delay caused by the path-length increase is negligible compared with the slowing effect itself (Jehle 2009). Pressure, humidity and temperature changes in the troposphere (the lowest ~ 11 km of the atmosphere) are the primary cause of refraction at X-band (Jehle 2009; Klawitter 2000), contributing typically ~ 50 times more than the ionosphere to the total delay. The stratosphere contributes only a negligible amount. The tropospheric delay is composed of the hydrostatic delay (associated with the “dry” troposphere) and the wet delay. The modeled hydrostatic component can be considered accurate on the order of ~ 1 mm (Bevis et al. 1996), while the wet component is associated with a higher uncertainty of maximum ~ 5 cm, according to calculations made during the preparation of Jehle et al. (2008) (the “wet” component, when modeled using the method employed here, deviates from ref-

erence values—generated by a millimeter-accurate ray-tracing model—with a standard deviation of ~ 4.5 cm). It should be added that the model input parameters pressure, temperature and humidity all influence the path delay uncertainty; the values stated above are valid for measurements made by the national weather service at a station near the imaged location, within a couple of hours of its acquisition).

The TSX product annotations include an approximate PD value, calculated using a simple atmospheric model that takes only the mean imaged terrain height and the viewing geometry into account (Fritz et al. 2008a). The model neglects real day-to-day atmospheric variations. While the annotated PD estimate yields much better range estimates than *without* their inclusion, they may nonetheless deviate from the true PD by approximately half a meter in the case of shallower viewing angles.

The atmospheric model used in this study to generate improved PD estimates is described in Jehle et al. (2008) and Collins and Langley (1996). It is a purely height-dependent model calibrated using ground-based measurements of the air pressure, temperature and water vapor. For each individual data take, the meteorological measurements were obtained from the weather station Payerne, about 6 km north of the test site. Given these values, and under the assumption of a “mean atmosphere” for mid-latitudes, the PD was calculated for each individual TSX data set and illumination geometry. This estimate, independent of the radar data, was then used to adjust the range timing values, effectively “removing” the PD from the range measurements.

3.2 GPS reference frame drift

The tests of geolocation accuracy are based on comparisons between imaged CR positions and their surveyed DGPS

Fig. 6 Relationship between the Swiss local geodetic reference system CHTRS95 and ITRS (modified from [Schneider et al. 2001](#)). As indicated by the connection between ITRS and ETRS89 in 1989, this is when the two reference systems were last in perfect correspondence

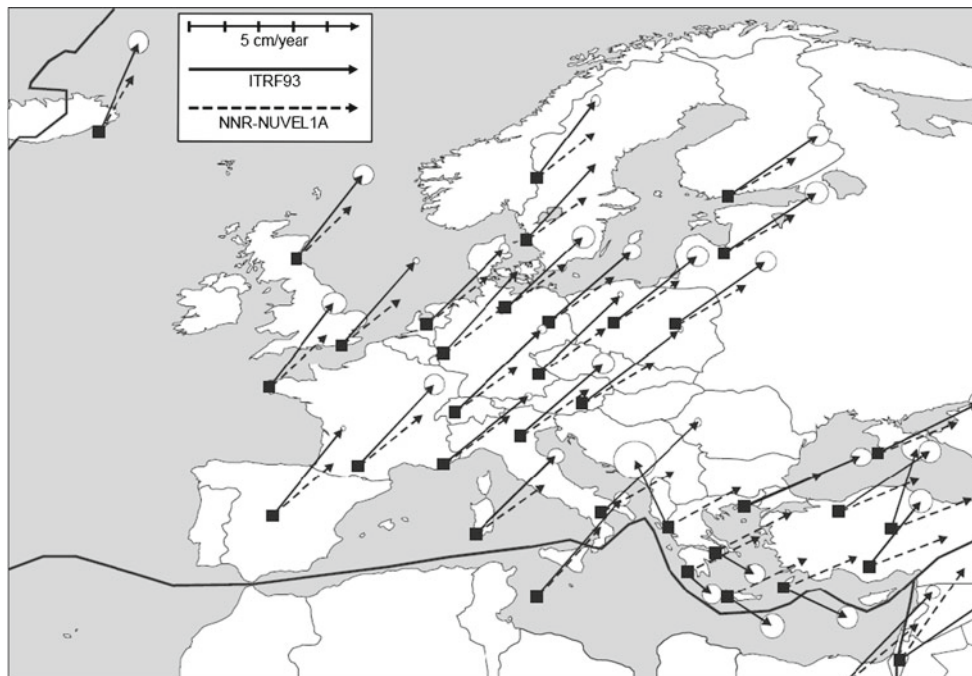
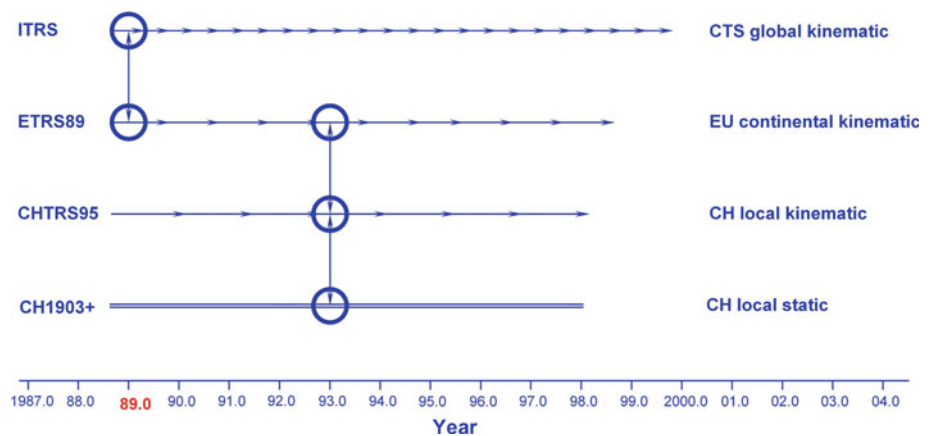


Fig. 7 Velocities of European reference stations within ITRS (from [Schneider et al. 2001](#))

positions. The comparison carries with it the assumption that the geodetic reference frames defined for the satellite positions and the CR positions are identical. In fact, this is often not the case. The reference frame used for the orbit state vectors is WGS84-G1150 (an instantiation of the World Geodetic System defined by the National Imagery and Mapping Agency, [NIMA 2004](#)) according to the TSX product annotations (format described in [Fritz 2007](#)). WGS84-G1150 and the International Terrestrial Reference Frame (ITRF) are identical to within approximately several centimeters at most, according to [NIMA \(2004\)](#) and [Schneider et al. \(2001\)](#) (although no documentation of the relationship is known to be publicly available). DGPS measurements of the CRs were provided in the Swiss Terrestrial Reference Frame CHTRF95, which is tied closely to the European

Terrestrial Reference System ETRS89 and is identical to it at the epoch 1993.0. Both are coupled to the stable part of the Eurasian continental plate. ETRS89, in turn, was defined in 1989 as identical to ITRS. Since 1989, these two frames have been slowly drifting apart due to their connections to different continental plates. Figure 6 illustrates the relationship between the reference systems. The transformations between the ITRS/ETRS/CHTRS systems are documented, for example in [Swisstopo \(2006\)](#) and [Boucher and Altamimi \(2001\)](#). The relative drift within ITRS is shown in Fig. 7, where Switzerland can be seen to be moving towards the north-east at a rate of ~ 2.5 cm/year. More precisely, the NNR-NUVEL1A model yields a motion of the Eurasian plate of 2.44 cm/year (calculated for example according to [Deutsches Geodätisches Forschungsinstitut 2011](#)) and a heading of 34.9° , with

a 90% confidence interval estimated empirically as better than 0.3 cm/year by Larson et al. (1997). The Eurasian plate movement was modeled using the NNR–NUVEL1A kinematic model (a version of the NUVEL1A model with “no net rotation”, or NNR), defined in DeMets et al. (1994). Thus, the total drift of CHTRS95 relative to ITRS between 1989 and the TSX acquisition period is ~ 50 cm.

By modeling the relative drift between CHTRS95 and ITRS for a given product acquisition date and projecting the movement into slant range and azimuth coordinates, the surveyed DGPS coordinates were transformed into ITRS. This was a necessary step for the subsequent geolocation tests and was found to be the primary source of residual errors in our previous study (Schubert et al. 2010), particularly in the along-track direction.

3.3 Solid Earth tide

Periodic deformations of the Earth cause vertical peak-to-peak fluctuations of up to ~ 40 cm at mid-latitudes. Relative to the tide-free state, this figure translates to SAR slant-range deviations of up to ~ 15 cm (i.e. ~ 30 cm peak-to-peak) and much smaller horizontal displacements of a few centimeter in the azimuth dimension. The SET model used is based on a computer program by Milbert (2011), an implementation of the model described in McCarthy and Petit (2004). The authors state the model to be accurate to ~ 1 mm. We compared the model calculations with the results of another model with a claimed accuracy of ~ 1 cm and observed agreement at the centimeter level, thus validating the tidal model used here to at least ~ 1 cm accuracy.

The tidal movement in 3-D of a single CR is shown for a full day (24.12.2010) in Fig. 8. Part (a) shows the movement separated into its vertical, north and east components, while (b) shows the same movement viewed from above. The dominant component is vertical, with a maximum deviation of about 16 cm shortly before 09:00 UTC. However, the acquisition on 24.12.2010 was at 17:25 UTC, when the vertical deviation was only ~ 12.2 cm; this translates to ~ 9 cm in the slant range direction for the corresponding ascending geometry.

In Fig. 9, the horizontal map positions (easting and northing) of one CR at the test site are plotted for the entire test period. The noisy orange data series includes all solid Earth effects, reflecting the irregular horizontal motion of the Earth's surface under the influence of the moon and sun. The blue series is obtained after removal of the modeled SET. It begins at the bottom left on 16.08.2009 and progresses in time towards the north-east. What remains is the movement of the CHTRF95 coordinates within the global ITRF reference frame due to plate tectonics; it is equivalent to the dotted arrow extending from Switzerland in Fig. 7. The total distance traveled over the test period of 484 days was 3.36 cm,

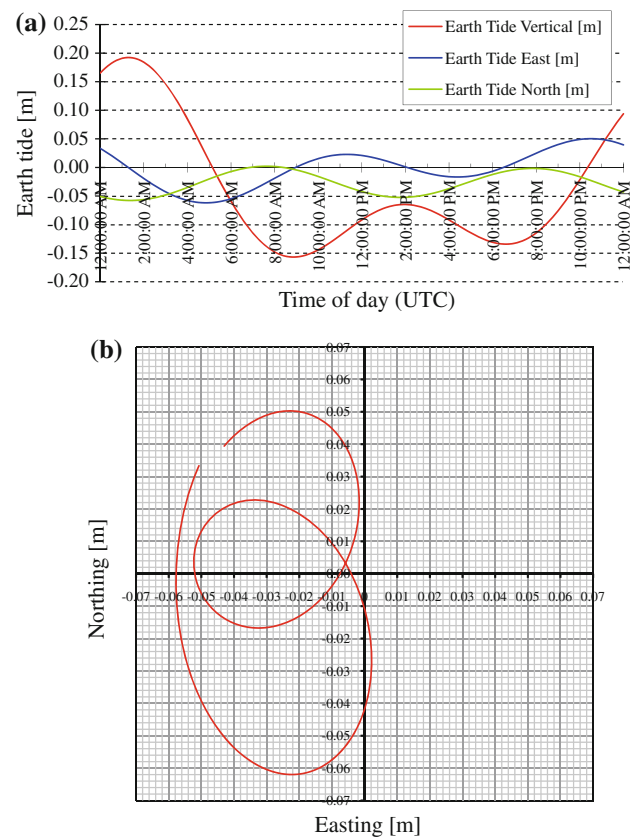


Fig. 8 Motion of a CR due to solid Earth tides over a 24-h period on 24.12.2010. **a** Variation of vertical, easting, and northing components of motion, and **b** easting and northing components of motion, top view

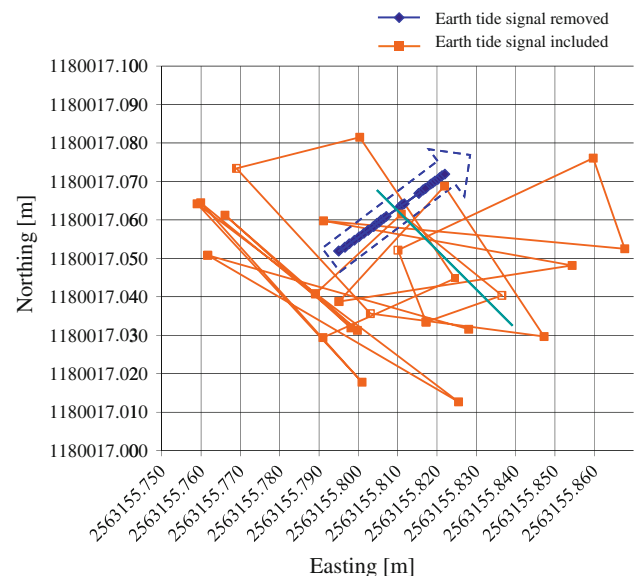


Fig. 9 Decimeter-scale horizontal motion of a GPS receiver over the 16-month study period. The orange points are affected by Earth tides as well as plate movement, while the blue points have had Earth tides removed, isolating the north-east drift of the Eurasian plate relative to the ITRF frame

corresponding to a 2.53 cm/year drift rate and a heading of 36.5° . The drift rate is well within the 90% confidence interval for the NNR–NUVEL1A plate model (see the previous section). Its heading is 36.5° , which compares very well to the 34.9° heading produced by the NNR–NUVEL1A model. We can therefore conclude that the observed drift is mainly due to the relative drift of the local and international reference frames.

4 Results

The core of the geolocation accuracy tests consists of a comparison between the imaged CR locations and their predicted positions, as described in Sect. 2.3. Four sets of test results are presented in the following sections, each subsequent result set incorporating an additional corrective measure. The results are presented in a scatter plot format, where the difference between the predicted and measured range and azimuth coordinates is plotted for each CR and each product. Two data points therefore exist for each TSX product. Different symbols distinguish ascending- from descending-orbit acquisitions. This distinction makes azimuth effects easily visible, as azimuth shifts have different signs for ascending and descending orbits. “Perfect” geolocation would result in range and azimuth differences of zero, which is at the center of the scatter plots in Figs. 10, 11 and 12. These plots extend from -1 to $+1$ m in both dimensions to simplify direct comparisons.

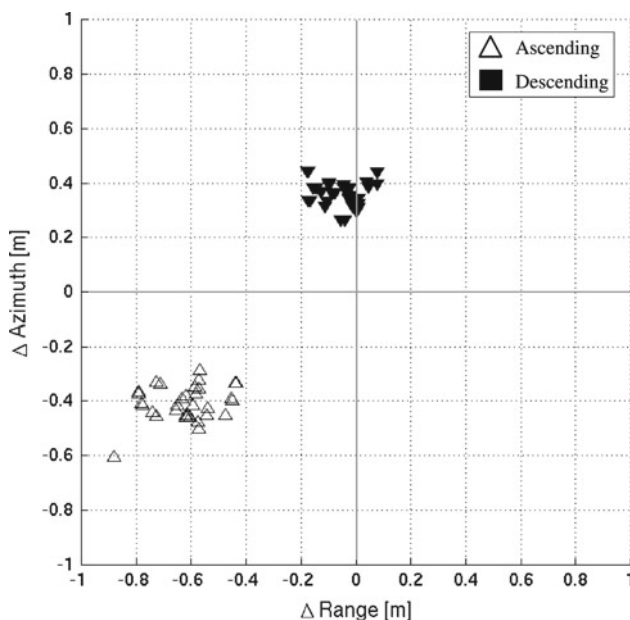


Fig. 10 Offsets between measured and predicted corner reflector positions using delivered product annotations (“out-of-the-box” test). Corrections applied: constant total atmospheric path delay delivered with the product annotations used to adjust radar timing

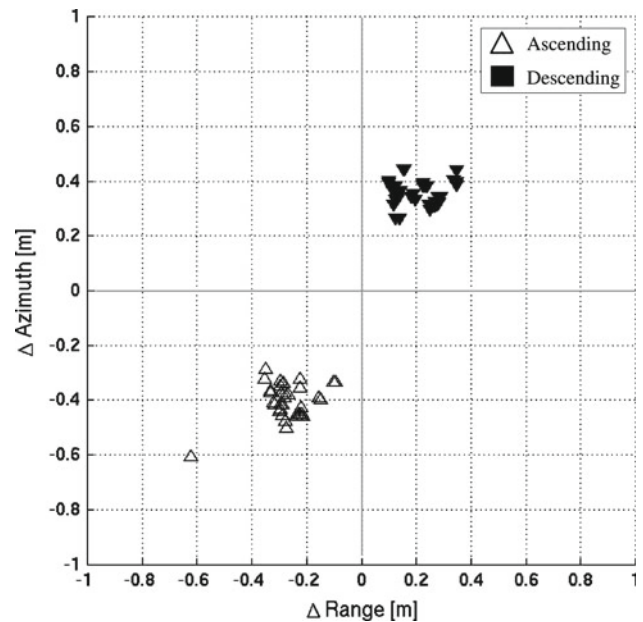


Fig. 11 Offsets between measured and predicted corner reflector positions. Corrections applied: constant path delay estimated by site- and time-specific atmospheric modeling, used to adjust radar timing

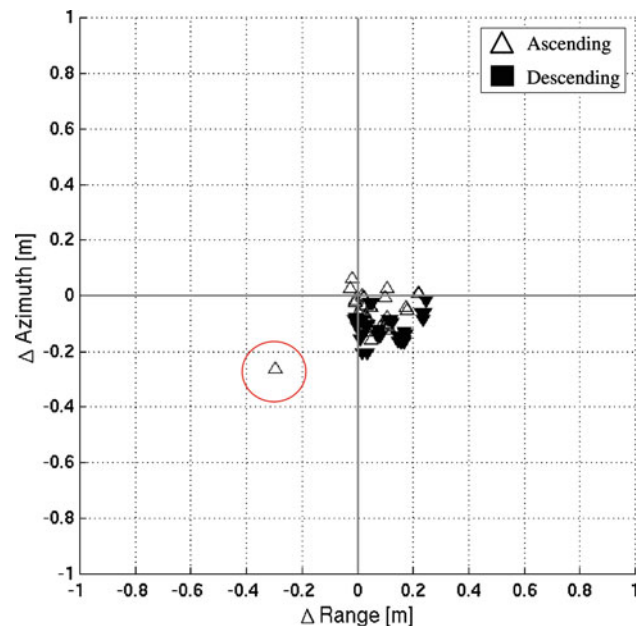


Fig. 12 Offsets between measured and predicted corner reflector positions. Additional correction applied: DGPS positions translated from their native CHTRS reference frame into the ITRS frame. The outlying point is from a partially snow-filled CR and can be discarded

4.1 “Out-of-the-box” accuracy

Many users of TSX image products will not be able to make the model-based corrections (PD and SET). Users not requiring better than ~ 0.5 m accuracy do not need to be concerned with these corrections. However, to achieve ~ 0.5 m

geolocation accuracy, one correction must be applied at minimum: the PD supplied in the delivered annotations should be subtracted from the range timing values. Applying that correction before geolocation tests resulted in the error scatter plot shown in Fig. 10. It is called the “out-of-the-box” result because of its sole dependence on the delivered annotations. Several points are worth noting about this result:

- The DGPS coordinates have not yet been projected into the ITRS; the separation between the ascending and descending point clouds is therefore not indicative of a geolocation error in the products themselves, but a measurement artifact (which will be corrected in a subsequent step).
- For reasons that will be discussed in Sect. 4.6, only products that were processed using version 4.5 (denoted v4.5) of the TSX processor are shown. This includes all products delivered on 27.03.2010 and later, as well as six earlier acquisitions that were later reprocessed using the newer version.

4.2 Correction for atmospheric path delay

The “out-of-the-box” test made use of the nominal PD, which is not linked to the true local weather conditions that existed at acquisition time. Replacing the nominal PD with a more refined model (see Sect. 3.1) results in a new error scatter plot, shown in Fig. 11. As the PD is purely a range effect, the data points have been shifted along the horizontal (range) axis compared with Fig. 10. The two main ascending and descending point clouds are now centered approximately on (0, 0).

4.3 Correction for GPS reference frame drift

The most dramatic improvement to the geolocation tests arises when the GPS coordinates are transformed into the ITRS reference frame before the CR predictions are made, in effect accounting for plate tectonics. Figure 12 shows the new scatter plot. In comparison to Fig. 11, the ascending and descending point clouds have merged, with a mean error now visibly smaller than 20 cm in both dimensions.

A notable outlier can be seen in Fig. 12 to the left of the main error cloud. The corresponding CR was from an ascending product from 02.12.2010. Heavy snowfall was observed the day before, confirmed by meteorological reports (from the Federal Office of Meteorology and Climatology, MeteoSwiss), with 10–20 cm of accumulation. It is therefore very likely that one or both of the CRs were partially filled with snow, altering the CR’s backscattering characteristics. In addition to this, the X-band radiation is partially scattered by snow, lowering the contrast between the CR backscatter and the surrounding snow-covered concrete. The difference

between a summer acquisition and the 02.12.2010 acquisition can be seen in Fig. 13. The snow-free reference is Fig. 13a; part (b) is the same extract from the 02.12.2010 product. Although both CRs are visible in Fig. 13b, they are less distinct; the arms of the “cross” shape have disappeared. Also, the backscatter from the surrounding background (clutter) has increased due to the snow cover. The outlying error estimate in Fig. 12 corresponds to the rightmost CR in Fig. 13b, whose predicted position is visibly offset from the center of the CR. We conclude that for this CR, the maximum backscatter does not correspond to its vertex position. This hypothesis is supported by the extremely low SCR of 9.3, which is ~ 20 times lower than for CRs in the other ascending products. We are, therefore, confident that this point can be discarded.

4.4 Correction for SET

The final correction applied to the CR positions is the correction for the SET. As seen in Sect. 3.3, the fluctuations are generally on the order of several centimeters, and they can be easily modeled and projected into the range and azimuth dimensions. Applying the corrections for SET perturbations, the error points converge even further; the result is shown in Fig. 14.

At this point, all major corrections have been made to the CR positions. Discounting possible residual errors inherent within the atmospheric and tidal models themselves, the remaining offsets must be due to a combination of measurement errors, centimeter-level effects (such as ocean tide loading), minor geometric calibration errors and residual uncertainties in the orbital state vectors. Excluding the outlier (snow-filled CR), the mean range error is +13 cm with a standard deviation of 3.3 cm; the azimuth error is -8.3 ± 4.4 cm. Given the various remaining error sources (the measurement of the CR phase center alone cannot be claimed to be more accurate than ~ 1 –2 cm), this represents a remarkably high sensitivity and calibration precision for a spaceborne SAR sensor.

The 3.3 cm standard deviation value in range is consistent with the observed ranging spreads of 2.6–5.9 cm estimated by Eineder et al. (2011) in a similar study (which also corrected for solid Earth tidal movements). Given the previously cited uncertainties associated with the atmospheric, tidal and tectonic models (~ 5 , ~ 0.3 and ~ 1 cm, respectively), it is remarkable that the total standard deviations in both dimensions of the corrected offsets is not > 3 –4 cm. Separating the dimensions: since the range spread should mainly be affected by path delay modeling errors, the ~ 3 cm standard deviation can be considered very small. The centimeter-level uncertainty in the solid Earth model results imply that the ~ 4 cm spread in azimuth is mainly due to orbital and timing errors augmented by solid Earth modeling errors.

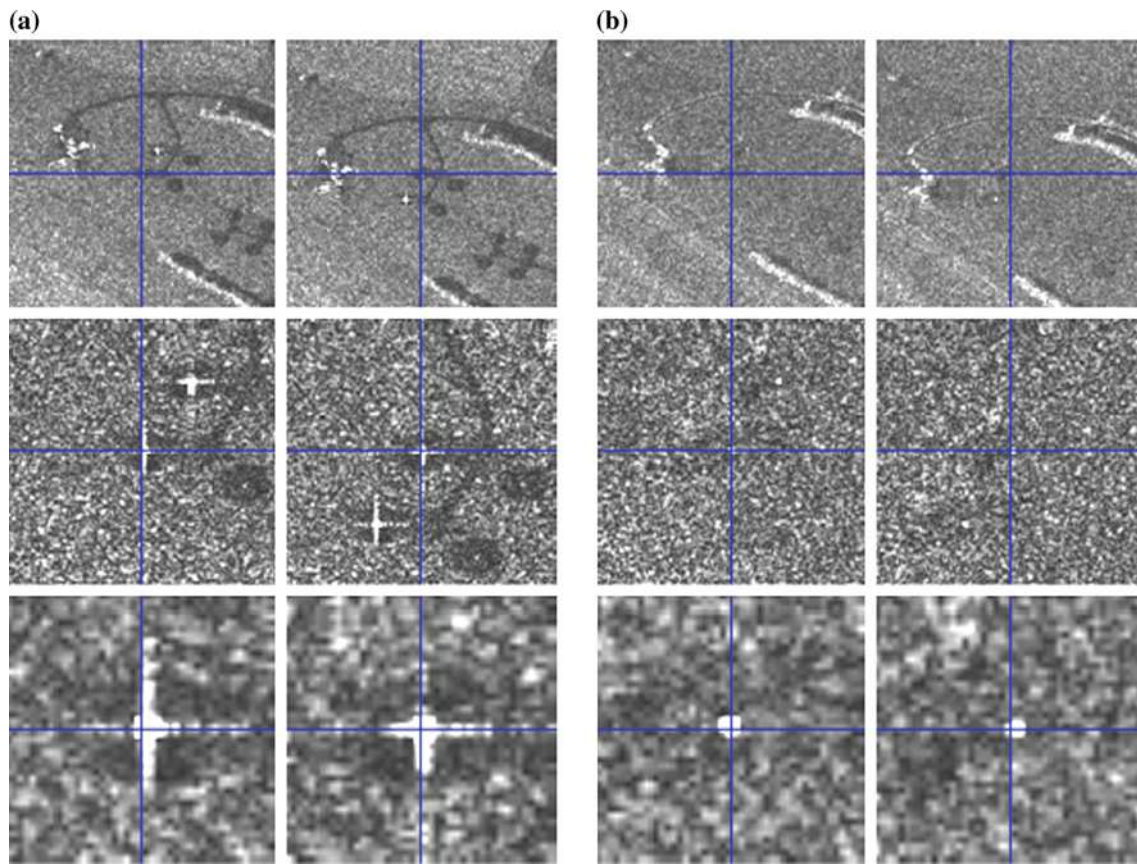


Fig. 13 Comparison of CR detection in snow-free and snow-covered scenes from the ascending orbit. The *crosshairs* represent predicted locations of the CRs. **a** Snow-free reference from 25.08.2010; SCR

(left) = 199.1, SCR (right) = 212.3. **b** Image from 02.12.2010 acquired after heavy snowfall; SCR (left) = 29.3, SCR (right) = 9.3

The error cloud is so compact that its offset from (0, 0) is probably due to a slight miscalibration of the sensor's timing constants. Since this is a simple matter of processor (software) calibration, it is not a flaw in the sensor design itself.

4.5 Further investigations: signal-to-clutter ratio, seasonal dependence and SAR processor version

Figure 14 was investigated further for significant correlation with the SCR (reliability of the imaged CR position) and time of year (seasonal reliability of the atmospheric model used to estimate PD).

To approach these two questions simultaneously, we plotted Fig. 14 again, encoding the SCR and seasonal information by varying marker color and size, respectively. The result is shown in Fig. 15a: SCR is color-coded with the marker size indicating one of three seasonal groups. Figure 15b is the same as (a), except that all points with SCR < 150 have been eliminated. Based on these plots, we conclude that

1. no significant correlation between season and the accuracy of the modeled PD is evident;

2. while extremely low SCR (i.e. <50) may occasionally indicate a false measurement of the CR location, the CR image analysis is usually robust enough to provide the correct location.

Concerning the processor version, the series of scatter plots presented in Sect. 4.1–4.5 did not include many products processed prior to 27.03.2010. This is because a newly calibrated TSX processor was used from that point onwards, v4.5. A comparison between error plots for v4.4 and v4.5 for six acquisitions processed twice, once for each processor version, is given in Fig. 16. Part (a) shows the v4.4 plot, (b) the v4.5 plot. Both scatter plots include correction for PD, plate tectonics and the SET, so they are directly comparable to Fig. 14. The improved v4.5 azimuth and range calibration is clearly evident.

In summary, the scatter plot obtained in Fig. 15a therefore represents our best estimate of the “inherent” TSX geolocation accuracy at the time of this writing. It testifies not only to TSX's extremely high geolocation accuracy, but also its long-term stability.

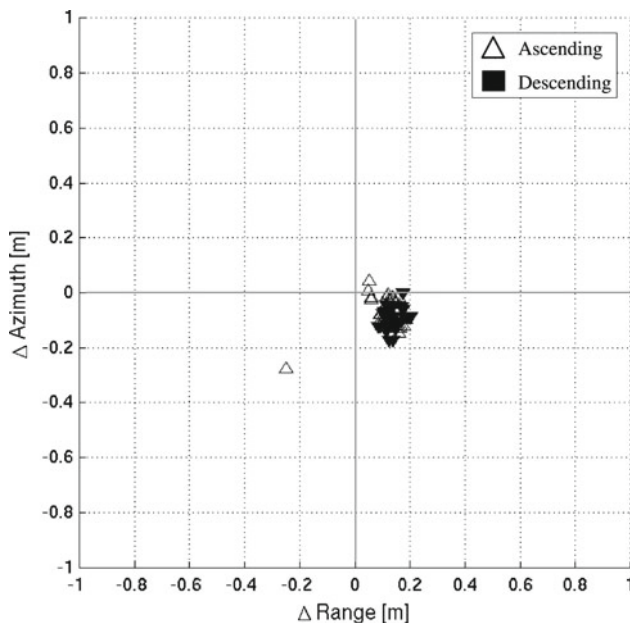


Fig. 14 Offsets between measured and predicted corner reflector positions. Additional correction applied: compensated local movements caused by solid Earth tides. The outlier to the left of the point cloud can be discarded as the reflector was partially snow-filled on that date, interfering with its nominal reflectance properties

4.6 Summary of geolocation accuracy estimates

The mean and standard deviation of the error scatter plots are listed in Table 1. The best estimate of TSX's current geolocation accuracy is *italicized* at the bottom of the table: under “typical” viewing conditions, the *consistency* (i.e. standard deviation) of the geolocation can be said to be within 3–4 cm, while the *absolute* location accuracy is on the order of a decimeter. The *absolute* accuracy could be improved by a further adjustment of the TSX processor's calibration constants, although it is already far below the 1-m requirement defined for TSX (Fritz and Eineder 2008b).

5 Conclusion

In this study, we acquired a 16-month time-series of high-resolution spotlight images from the TSX sensor over a test site in Switzerland. Four trihedral CRs were deployed and their positions were surveyed with approximately centimeter accuracy. We compared the measured CR positions with their imaged positions, based on the delivered product annotations and corrections made for atmospheric path delay, plate tectonics and SETs. The resulting error distribution is very compact and allows us to assign the TSX system an *absolute* geolocation accuracy of approximately 13 cm in range and 7 cm in azimuth, with a standard deviation of 3–4 cm in both dimensions.

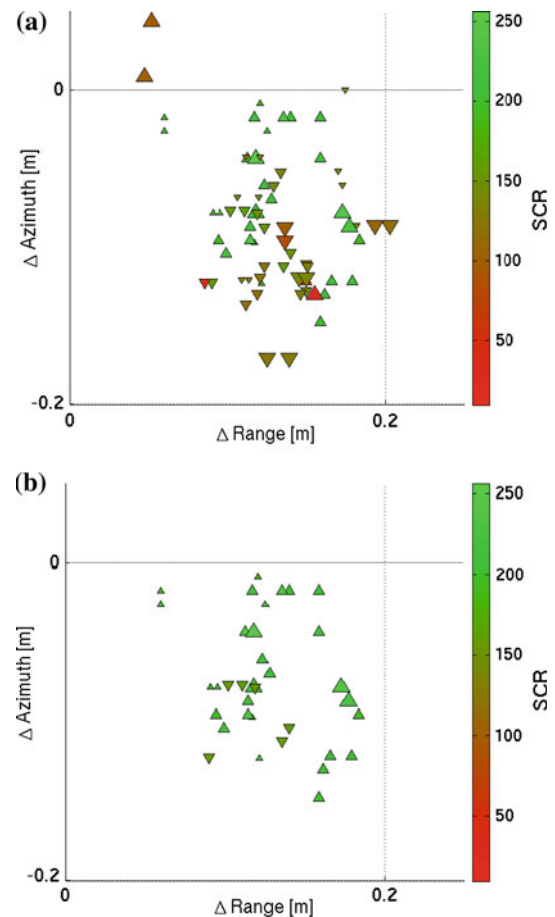


Fig. 15 Offsets between measured and predicted corner reflector positions, color-coded according to the image signal-to-clutter ratio, and marker sizes according to season. *Large triangles* are from winter acquisitions, *medium triangles* from spring and autumn, and *small triangles* from summer acquisitions. **a** Errors for all products and **b** errors after removal of points with $SCR < 150$

The sensor has proven to be sensitive enough to (incoherently) detect solid Earth movements on the order of several centimeters. At the same time, this study provided indirect evidence for the validity of the path delay and solid Earth tidal and tectonic models used during the study. Further work would be necessary to validate the models used in non-temperate latitudes, e.g. high arctic and equatorial regions.

Due to ever-improving data availability from Earth-observation platforms, there is a current push towards an increased utilization of image time-series from one or more sensors. TSX has set a new geolocation standard for spaceborne SAR systems. However, this accuracy would be severely compromised were the atmospheric path delay not systematically mitigated. Therefore, we highly recommend the widespread inclusion of an atmospheric model in the product annotations, calibrated for the acquired scene by taking the best

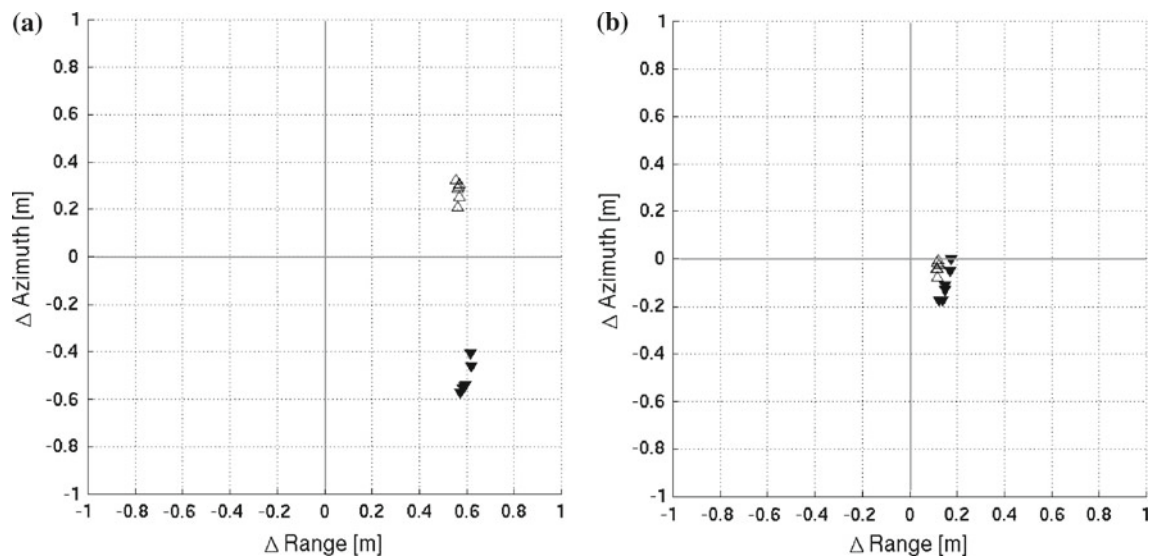


Fig. 16 Comparison of offsets between measured and predicted corner reflector positions for products processed using the v4.4 or v4.5 processor. Scene-specific path delay, tectonic, and tidal corrections have all

been applied. **a** Offsets for the six v4.4 products and **b** offsets for the same six acquisitions as **a**, but processed using the v4.5 software

Table 1 Offset statistics according to corrective measures implemented (all using v4.5 processor)

Corrective measures	Dimension	Ascending (m)	Descending (m)	Total (m)
“Out-of-the-box” path delay	Range	-0.61 ± 0.10	-0.05 ± 0.07	-0.32 ± 0.30
	Azimuth	-0.40 ± 0.05	0.36 ± 0.04	-0.00 ± 0.38
Scene-specific path delay	Range	-0.26 ± 0.06	0.20 ± 0.08	-0.02 ± 0.24
	Azimuth	-0.40 ± 0.05	0.36 ± 0.04	0.00 ± 0.38
+ Plate tectonics	Range	0.06 ± 0.07	0.09 ± 0.08	0.08 ± 0.07
	Azimuth	-0.06 ± 0.05	-0.12 ± 0.05	-0.09 ± 0.06
+ Earth tides	Range	0.13 ± 0.04	0.14 ± 0.03	0.13 ± 0.03
	Azimuth	-0.07 ± 0.05	-0.10 ± 0.04	-0.08 ± 0.04

The number of points contributing to the statistics = 69

available meteorological measurements and topography into account.

Acknowledgments The authors would like to thank the German Aerospace Center (DLR) for acquiring and providing the TerraSAR-X products, EADS/Astrium for lending us trihedral corner reflectors, and the Swiss Air Force for permission to place corner reflectors near Torny-Le-Grand. We are also grateful to Swisstopo for providing the ITRF coordinates and velocities of the DGPS reference stations, as well as the CR deployment and surveying team for their tireless dedication. Finally, we thank the four reviewers who suggested useful improvements for the final version of the paper.

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